

Contribution of CH₄ to Multi-Gas Reduction Targets

The Impact of Atmospheric Chemistry on GWPs

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Abstract: Weighting indices can be used to express emissions of different greenhouse gases (GHGs) in CO₂-equivalent terms. The choice of a specific weighting index is often thought to determine the emphasis on non-CO₂ GHGs in a multi-gas target. Although the choice of a 100-year Global Warming Potential (GWP) for the Kyoto Protocol narrows the range of possible index values, we use the example of CH₄ to show that this choice does not constrain the GWP to a single number. Projected changes in atmospheric CH₄ concentrations due to interactions with CO, NO_x, and OH produce direct GWPs ranging from 12.5-15.5. This increases to 14-24 when the range of uncertainty in indirect effects due to tropospheric ozone (O₃) and stratospheric water vapour is included. We examine the impact of the calculated range in GWPs on the optimal contribution of CH₄ to U.S. emission reductions up to 1.0 GtC_{eq}. Our results show that for a 100-year GWP there is still significant variation in effective CH₄ reductions for a multi-gas target. This has implications for the overall contribution of CH₄ and other non-CO₂ GHGs to both short-term and long-term strategies focussed on the mitigation of climate change impacts.

1. INTRODUCTION

Abatement of non-CO₂ GHG emissions can increase flexibility and lower the costs of meeting emission targets relative to CO₂ abatement alone. For a multi-gas target, other gases are often expressed in terms of CO₂ equivalence

425

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through weighting indices. These indices are the primary way to avoid lengthy model calculations when determining the mix of GHG reductions required to meet a target such as that set by the Kyoto Protocol. However, it is widely recognised that weighting indices may misrepresent the contribution of non-CO₂ GHGs to reduction strategies, particularly when considering the impact of climate change, atmospheric chemistry, and trace gas emissions on methane (CH₄).

Methane is the second most important anthropogenic greenhouse gas, responsible for almost 20% of the increase in forcing since pre-industrial times (IPCC, 1996). The atmospheric lifetime of CH₄ is highly sensitive to emissions of trace gases such as CO, NO_x and NMHCs. In the future, CH₄ lifetime may also be affected by climate-change-induced increases in atmospheric temperature and water vapour that affect the production rate of OH, the primary sink of atmospheric CH₄ (Wuebbles *et al.* 1999).

CO₂-equivalent weighting indices for CH₄ are highly dependent on time horizon, due to the order-of-magnitude difference between CH₄ and CO₂ lifetimes. This produces a wide selection of weighting indices (Table 1), ranging from 65 for a time horizon of 20 years, down to <1 for an infinite time horizon over which biogenic CH₄ is completely removed and fossil CH₄ oxidised to CO₂. Attempts to consider this full range will obviously produce a large variation in the potential contribution of CH₄ to GHG reduction strategies. For this reason, policies select a specific type of weighting index and appropriate assumptions on which to base that factor, taking into account the type of issues addressed by that policy. For example, the Kyoto Protocol chose a 100-year GWP to calculate the CO₂-equivalence of non-CO₂ GHG emissions (UN, 1997). This subjective choice narrows the range of possible values for weighting indices, and constrains potential CH₄ contributions to emission targets. However, a recent revision in the radiative forcing of CO₂ (WMO, 1999) has raised the 100-year GWP for CH₄ from its value of 21 at the time of the Protocol up to 24. This revision illustrates the fact that GWPs are highly dependent on current scientific uncertainties, as well as on future projections of trace gas emissions and climate change.

Table 1. CH₄ weighting indices (direct effects only).

WEIGHTING INDEX			VALUE	REFERENCE(S)
Global Warming Potential	GWP pulse	20 years	37-53	IPCC 1995, 1996; Bruhl, 1993; Lelieveld <i>et al.</i> 1993; 1998
		100 years	12-21	
		500 years	4-7	
	GWP sustained	6-35	Fuglestedt <i>et al.</i> 1996	
Economic Damage Index	EDI		3-50	Hammit <i>et al.</i> 1996
Scenario-Based Index	SBI		6-30	Kandlikar 1995
Forcing Equivalence Index	FEI	20 years	39-45	Wigley 1998
		100 years	20-28	

Ignorance of the fact that the choice of one value for a given weighting index and time horizon may reflect neither scientific uncertainties, nor its dependence on future atmospheric conditions, tends to produce a false sense of certainty. In this study we illustrate the extent to which the choice of one weighting index, particularly one as dependent on future conditions and current uncertainties as a GWP, still allows for a significant range in values. We then assess the impact of this range on the effective contribution of CH₄ to a range of GHG emission reductions in 2010.

2. GWPS: RELATING CH₄ TO CO₂-EQ. EMISSIONS

Global warming potentials for CH₄ emissions are determined by the relationships between CH₄ and CO₂ concentrations and radiative forcing. Each of these relationships in turn is affected by a number of scientific uncertainties and time-dependent factors (*Fig. 1*). Relating CH₄ emissions to radiative forcing, then to CO₂-equivalent emissions, depends on factors such as the impact of trace gas emissions and climate change on the lifetime of CH₄; absorption efficiency; the magnitude of indirect radiative effects from tropospheric O₃, CO₂, and stratospheric water vapour formation; spatial and temporal forcing distributions; and our understanding of the carbon cycle.

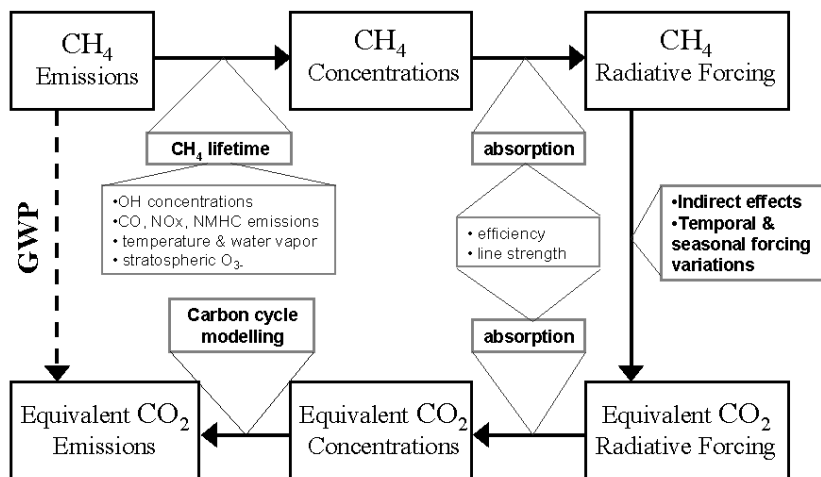


Figure 1. Primary variables affecting the chain of relationships in converting CH₄ emissions to their CO₂ equivalents. In actual GWP calculations, CH₄ radiative forcing is converted directly to CO₂-equivalent concentrations using the forcing relationship given in IPCC (1990).

Changes in each of these relationships arise from two sources: improvements in scientific understanding, and time-dependent changes in atmospheric temperature and composition. Limits to scientific understanding are addressed by incorporating reasonable uncertainty ranges, recognising that certainty may improve over time. This is crucial for model-dependent factors such as the indirect effects of CH₄ on tropospheric O₃ (IPCC, 1995). However, the recent 15% decrease in CO₂ radiative forcing reported by WMO (1999) highlights the need to consider uncertainty in *all* factors.

When relating CH₄ emissions to radiative forcing, the lifetime of CH₄ is the factor most sensitive to time-dependent changes. CH₄ lifetime is mainly determined by atmospheric OH levels. OH is removed by CH₄ and its oxidation product, CO. In the CH₄-OH-CO cycle, increasing methane emissions also increase atmospheric CO, enhancing OH removal, slowing down CH₄ oxidation, and allowing CH₄ concentrations to build up in the atmosphere (Wuebbles *et al.* 1999). Due to this cycle, methane is removed by OH on an ‘adjustment’ time scale that exceeds the chemical lifetime by approximately three years (e.g., Prather, 1994; IPCC, 1996).

OH production and consumption is also affected by NO_x and NMHCs emissions, stratospheric O₃ concentrations, atmospheric temperature, and water vapour levels. Although current changes in OH production appear to be balanced by consumption at this time (Prinn *et al.* 1995), increasing trace gas emissions are expected to decrease OH levels in the future (e.g., Thompson, 1992; Brasseur *et al.* 1998; Lelieveld *et al.* 1998).

Comparison of the radiative forcing produced by CH₄ emissions with that of CO₂ is additionally uncertain because of indirect forcing effects that result from the production of tropospheric O₃, stratospheric water vapour and CO₂ through the oxidation of CH₄. Estimates of the first two indirect effects range from 10-30% of the total GWP for CH₄ (e.g., IPCC, 1995; Bruhl, 1993; Lelieveld *et al.* 1993, 1998; Fuglestedt *et al.* 1996).

In the next sections, we explore the dependence of the 100-year GWP for CH₄ on atmospheric chemistry. We then evaluate the impact of the uncertainty induced by future projections of trace gas emissions and indirect effects on the optimal contribution of CH₄ to GHG emission reductions, and the cost reduction obtained by including CH₄ in a CO₂ reduction strategy.

3. IMPACT OF TRACE GAS EMISSIONS ON GWPS

We use the Integrated Science Assessment Model (Jain *et al.* 1994, 1999) to calculate the direct GWP of CH₄, taking into account the CH₄-OH-CO cycle and the impact of changes in trace gas emissions on CH₄ lifetime. Indirect effects are assumed to be 30% of the direct GWP (IPCC, 1995).

GWPs corresponding to six scenarios are shown in Table 2. Trace gas emissions are based on the IS92a 'business-as-usual' emission projections (IPCC, 1999), while CO₂ is held constant at 1990 levels.

NO_x emissions have already been identified as a key factor affecting the relationship between CH₄ emissions and concentrations (Ksheshgi *et al.* 1999). Comparison of GWPs for scenarios 2 and 6 extends this conclusion to show that changes in OH production are the most crucial factor in determining the sensitivity of the GWP of CH₄ to atmospheric composition for the scenarios examined. Although NO_x emissions are mainly responsible for the OH changes, OH may also be affected by climate change impacts unresolved by our model. The range in GWPs shown in Table 2 is likely to increase for a greater variety of emission and climate change scenarios.

Table 2. Impact of trace gas emission scenarios on the 100-year GWP for CH₄. Indirect effects are 30% of the direct GWP. Projections are given by the IS92a scenario (IPCC, 1999).

SCENARIO	GWP	
	Direct	Total
1. Constant CH ₄ concentrations (1990 levels)	14.6	19.0 (16-23)
2. Changing CH ₄ and CO emissions, constant NO _x emissions (1990)	12.5	16.3 (14-19)
3. Changing CH ₄ emissions, constant CO and NO _x emissions (1990)	12.9	16.8 (15-20)
4. Changing CH ₄ concentrations (regardless of other emissions)	13.5	17.5 (15-21)
5. Changing CH ₄ and NO _x emissions, constant CO emissions (1990)	14.3	18.6 (16-22)
6. Changing CH ₄ , CO and NO _x emissions + additional 10% increase in OH production due to increasing temperatures and water vapour	15.5	20.1 (17-24)

4. IMPACT OF GWPS ON CH₄ ABATEMENT

Future projections of trace gas emissions examined in these 6 scenarios produce a total (direct + indirect) 100-year GWP of CH₄ ranging from 16.3 to 20.1. When the uncertainty in indirect effects is taken into account, this range increases to 14-24, implying that uncertainties in the indirect effect are more likely to increase than decrease the importance of CH₄ relative to CO₂.

We evaluate the impact of these ranges in terms of two measures of the potential for CH₄ to alleviate CO₂ reductions: (1) optimal CH₄ emission reductions (*Fig. 2*), and (2) the cost savings achieved by including CH₄ (*Fig. 3*). CH₄ mitigation cost estimates are those given by the U.S. EPA (1999) for reductions from coal mining, natural gas, landfills, and animal manure for the U.S. in 2010. CO₂ abatement costs for the U.S. are calculated by the SGM, a computable general equilibrium model designed to address issues associated with global change (Edmonds *et al.* 1998). Optimal partitioning between CH₄ and CO₂ reductions is calculated by equating marginal abatement costs (Hayhoe *et al.* 1999).

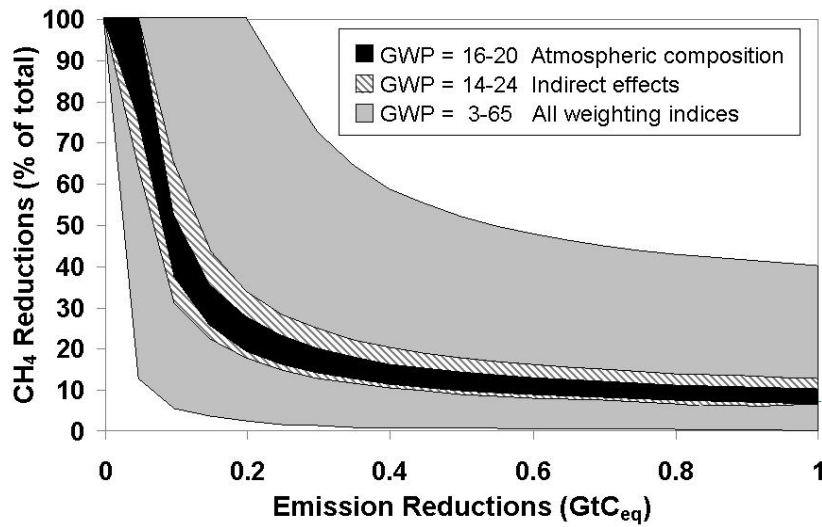


Figure 2. Optimal contribution of CH₄ to CO₂+CH₄ emission reductions, expressed as % of total CO₂ and CH₄ reductions. Costs are for U.S. emission reductions ranging from 0 to 1.0 GtC_{eq} in 2010. Three areas correspond to ranges in: (1) 100-year CH₄ GWP due to changes in atmospheric composition (black), (2) 100-year GWP due to uncertainties in indirect effects (striped), and (3) all weighting indices over time horizons from 20 to 500 years (grey).

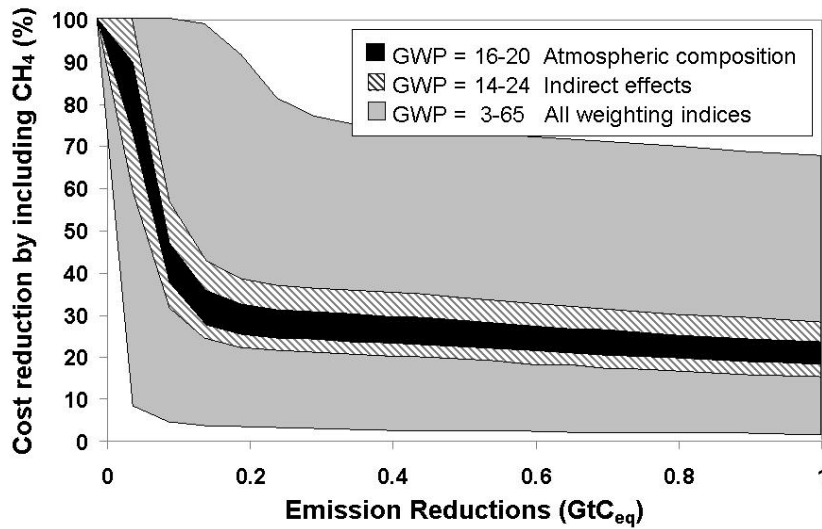


Figure 3. Cost saving achieved by including CH₄ in a CO₂-only reduction strategy, expressed as % relative to CO₂-only costs. Costs and coloured areas are as described in Fig. 2.

5. CONCLUSIONS

Weighting indices used to compare CO₂ with other GHG emissions are sensitive to a number of conditions that affect the emphasis on CH₄ and other non-CO₂ GHGs in multi-gas targets. Selection of a GWP with a time horizon of 100 years can eliminate approximately three-quarters of the sensitivity for CH₄. However, these results show that even with a specified time horizon, there is still a significant variation in GWP values.

Many factors affect the long chain of relationships between CH₄ and CO₂-equivalent emissions. We found that two of these factors, future changes in atmospheric composition and uncertainties in indirect radiative effects, can vary the optimal contribution of CH₄ to a joint CO₂/CH₄ target by at least 4-7%. The effect on the cost savings achieved by incorporating CH₄ into a CO₂ reduction scheme is even larger, ranging from 5-13%.

The impact of trace gas emissions on GWPs suggests that the choice of a specific GWP for CH₄ emission reductions is quite sensitive to projections of NO_x, CO, and other gases. Current understanding of tropospheric chemistry limits our ability to model CH₄, OH, and other chemically active gases, and hence our ability to determine weighting indices that capture the full effects of CH₄ on climate (Wuebbles, 1996).

The Kyoto Protocol does not include emission reductions for non-GHG gases that clearly affect the relationship between CH₄ emissions, concentrations and GWPs. As it is likely that emissions of these gases will be affected by controls on sources shared with CO₂ and other GHGs, non-GHG trace gases should be considered in future control strategies.

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